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**GEOLOGICAL-PALEOOCENOGRAPHIC ENVIRONMENTS OF THE LOWER MAYKOP DEPOSITS OF AZOV-KERCH REGION**

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**ГЕОЛОГО-ПАЛЕООКЕАНОГРАФІЧНІ УМОВИ ФОРМУВАННЯ НИЖНЬОМАЙКОПСЬКИХ ВІДКЛАДІВ АЗОВО-КЕРЧЕНСЬКОГО РЕГІОНУ**

Detailed investigations of the lithological structure of Early Oligocene (three types of Lower Maykop sections were distinguished) were made. The lithological-facial schemes of Lower Maykop in Azovo-Kerch region, which reflect the spatial-temporal features of strata forming components (terrigenous rocks and argillites) distribution in sections were built. Mineralogical-petrographic composition of the rocks was characterized. The features of the lithological structure of Lower Maykop deposits were studied. Geological-paleoceanographic model of Azov-Kerch region on the beginning of Early Oligocene was built.

*Keywords:* Maykop series, section types, lithological-facial scheme, lithites, paleoceanography.

Детально вивчено літологічну структуру товщі ранньоолігоценного віку (виокремлено три основних типи розрізу нижньомайкопських відкладів). Побудовано літолого-фаціальні схеми нижнього майкопу Азово-Керченського регіону, які відображають особливості просторово-вікового розподілу товщеформуючих компонентів (теригенні породи та аргіліти) у розрізі. Охарактеризовано мінерально-петрографічний склад досліджуваних порід. Вивчено особливості літологічної структури розрізів товщі ранньомайкопського віку. Побудовано геолого-палеоокеанографічну модель Азово-Керченського регіону на початок раннього олігоцену.

*Ключові слова:* майкопська серія, типи розрізу, літолого-фаціальна схема, літміти, палеоокеанографія.

**INTRODUCTION**

Deposits of Maykop age is an interesting subject in oil- and gas-bearing topic in the range of oil and gas region of the Southern Ukraine (Denega et al., 1998; Bohaec et al., 1986). Despite the long history of studies of these deposits, until now the paleoceanographic features of their accumulation, including conditions of the aleurite-psammitic formations, that are the potential reservoir-rocks for oil and gas, have not been sufficiently studied.

To resolve this problem, the available lithologic, mineralogical and petrographic, geological and geophysical materials on the Lower Maykop beds of the Azov-Black Sea area were generalized and reinterpreted.

**GEOLOGICAL SETTING**

In tectonic respect the studied area covers three depressions: Northern Azov, Indol-Kuban and Sorokin as well as the separating them mega-uplifts (Central-Azov, Crimean and Caucasian). In stratigraphic respect the studied deposits are represented by Lower- and Upper Planorbelova and Ostrakoda regional stages of Early Oligocene (the Lower Maykop beds).

The presently uncovered thickness of deposits in the region varies from the first metres (Plane Crimea) to 1500 m (Near-Kerch Black Sea shelf) (fig. 1). They are absent in mega-anticlinorium of the Mountain Crimea and vault areas of Central-

Crimean and Central-Azov uplifts.

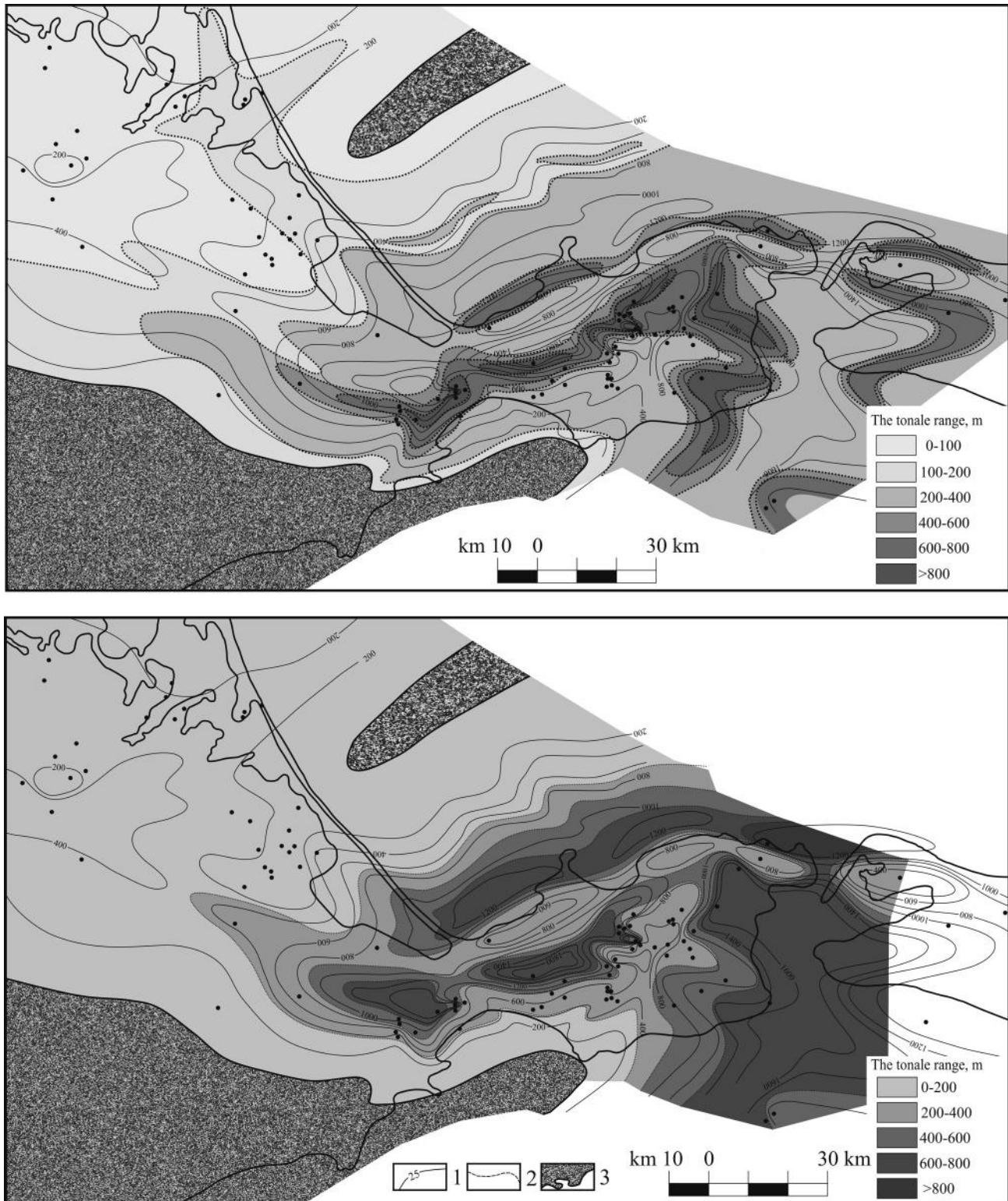
The Lower Maykop series of the Azov-Black Sea area is characterized by a rather monotonous lithology with the predominance of clayey rocks that vary by content of aleuro-psammitic and carbonate (sideritic) components, the number and species spectrum of faunal remains.

As to lithology there are three types of cross-sections of the Lower Maykop deposits: northern, central and southern.

The Southern type of cross-section is the most complete and thick. It is developed in Indol-Kuban and Sorokin depressions and is characterized by a predominance of psammitic-aleurolitic rocks in the lower part of the sequence and their absence in its top. The uncovered thickness reaches 1500 m, and the predicted one – more than 2000 meters (figs. 1, 2).

The central section type is traced in eastern part of Plane Crimea (such areas as Strylkova, Peredova, Jankoiska area, Chongarska etc.). It is characterized by the absence of basal layers of the planorbelitic horizon and extensive development of the psammitic-aleurolitic rocks on the top of the sequence (fig. 2). Uncovered thickness of deposits is not more than 450 m.

The Northern type of cross-section is located in the north-eastern parts of Crimea and in the northern area of near-Black Sea (Genicheska area etc.). It is represented by a top of Upper Planorbelova and

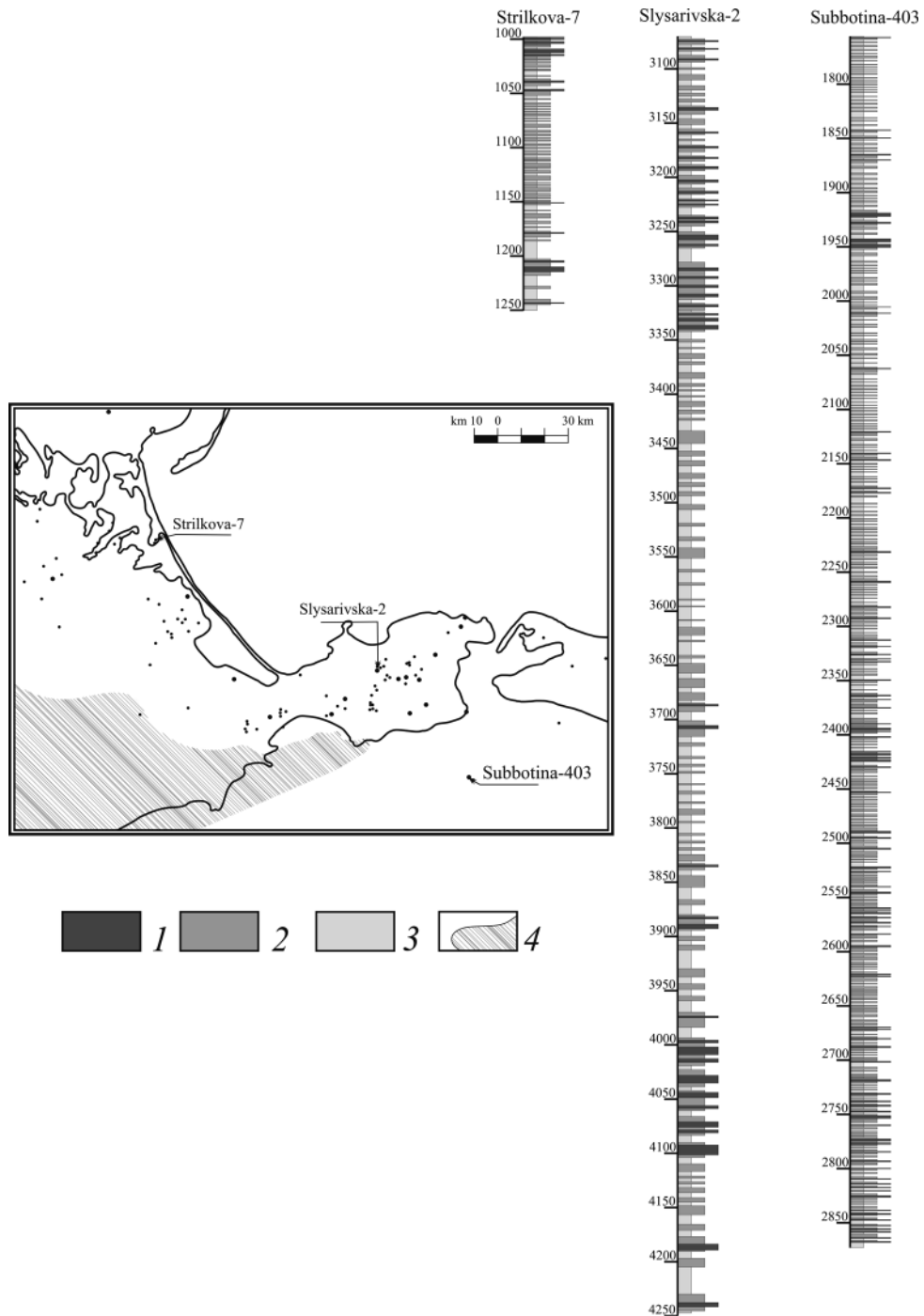


**Fig. 1.** Lithological-facial schemes (izolites: siltstones and sandstones (a), argillites (b)) of Lower Maykop of Azovo-Kerch region:

1 - isolithes; 2 - areas borders; 3 - deposits absent.

Ostrakoda horizons (thickness up to 200 m), and is characterized by a uniform distribution in cross-section of sandstones and siltstones (total thickness up to 40-70 m) (fig. 2).

Analysis of peculiarities of spacial-temporal development of the sequence forming components (sandstones and siltstones +argillites) in the Lower Maykop section allowed finding a number of peculiarities (fig. 1).



**Fig. 2.** Types of Lower Maykop sections in Azovo-Kerch region:  
 1 – sandstones; 2 – siltstones; 3 – argillites; 4 – deposits absent.

The lateral distribution of the total thickness of the sandstone-siltstone horizons is clearly correlating with the variations of its total capacity, showing the maximum values on the territory of Kerch Peninsula and in the Near-Kerch Black Sea shelf. Several fields of their maximum development have been localized.

A significant sublatitudinally elongated area with isolith values of 600 m is located in the the Kerch Peninsula (fig. 1). It extends from the Vlady-

slavivska area in the west to the Prydorozhna area in the east and further to Dubrovskia area in the south. The strip of smaller values of this parameter (300 m) borders the area. In addition, in this area the development of a number of small in size, highly elongated in the horizontal, sublatitudinal strike isolith areas with values of above 500 m are forecasted. These fields mostly tend to northern and southern slopes of syndepositional.

Within the northern territories of the region (eastern Plane Crimea and Near-Sivash Sea) local areal of maximum isolith values forms the area of submeridional stretch in the Strilkova with branches in Genichesk direction (north), Ust-Salhirska (south) and Jankoiska (west) structures. The maximum value of this parameter within it's ranges does not exceed 120-150 m and make 40 % of thickness of the total sequence. This area structurally tends to the axial part of the depression.

Probably quite a large areal of psammitic – aleurolitic rocks (total thickness of more than 500 m) of the Lower Maykop beds is localized in southern regions of the area (Subbotin area, East Black Sea trough, north slope of Shatsky high).

By the percentage in the sequence of the Lower Maykop of siltstone and sandstone horizons a stretched area of values of over 50 % is recorded. It stretches from south to north towards the prospects: Subbotina-Fontanivska-Strylkova. Its branches towards Vladyslavivska, Ust-Salhirska and Jankoiska prospects are observed.

#### DESCRIPTION OF MINERALOGICAL AND PETROGRAPHIC COMPOSITION

The Lower Maykop beds in the region, as it is noted above, represent rather uniform clay layers with different content of silt and carbonate-psammite (siderite) components, a number and range of faunal species remains. Intercalations of auleritic-psammitic varieties tend to lower or upper layers of the sequence. In the range of Indol-Kuban depression they form a specific basal Dyurmensk formation with total thickness of 20-650 m.

This transgressive formation rests on a highly dissected surface of Eocene-Palaeocene sediments and is composed of layers, lenses, packs of mudstones, siltstones, sandstones, rarely gravelite-breccias and of fine gravel conglomerates, with total thickness of 17-625 m. In the sequence, according to (Plachotny et al., 1971; Samarsky, 1983), from 2 to 7 lithological packages can be distinguished, which differ significantly by thickness, lithology and mineralogical and petrographic features in both the same and different structure. The most complete section of formation was found in Fontanivska, Hornostayevska, Bulganakhska and other areas.

*Sandstones* and siltstones of Dyurmensk formation are variegated, greenish-grey, light grey and dark grey, grey, brownish-grey, the colors are caused by mineral composition of cement and admixture of autigenous minerals (glauconite). The thickness

of individual psammitic layers usually does not exceed 20 m. The terrigenous deposits very often are facially replaced by clayey rocks.

The rocks are clayey, medium- and fine-grained, of olihomictic or more seldom of polymictic composition of chlorite-carbonate, hydromica, hydromica-chlorite, rarely silica-carbonate-chlorite and quartz cement of regenerative contact-interstitial and interstitial types. Their texture is layered, lense-layered or lense-spotted. Structure is psammitic.

Clastic material (70-90 %, 0,1-0,2 to 0,3-0,5 mm) is well sorted and rounded with grains of quartz (85-90 %) and feldspar (2 %), with the inclusion of fragments of siliceous rocks (1 %), light green orthorhombic pyroxene (hypersthene), scales of mica (1 %) and fragments of microfauna of carbonate composition. Authigenic minerals – glauconite (10-15 %), pyrite (3-5 %) and siderite. Fractures are filled by leicoxen or light bitumen.

*Siltstones* are clayey, sometimes sandy, grey, dark grey, greenish-grey with brownish tinge, aleuritic-psammitic structure or lense-layered texture. Clastic material the content of which reaches 70-75 %, ranging in size from 0,04 to 0,2 mm, of angular shape is unevenly distributed and represented by grains of quartz, feldspar and more seldom by grains of phosphate minerals and muscovite scales.

Accessory minerals are – garnet, zircon, epidote, authigenic – concretions of glauconite (20 %) and fine-grained siderite (20 %, often bordering grains of quartz). Carbonate content in rocks varies from 0 to 9,6 %, but the average is 2,5 %. The regular growth of carbonate content in rocks is noted from the northern to the southern areas of the region (table 1). Cement has hydromica, hydromica-chlorite, hydromica-galuazite rarely silica-carbonate-chlorite composition of contact-interstitial type.

*Argillites* are brownish-grey, dark grey to black, dense, heavy, sometimes lense-layered with layers of (1-5 mm) sandstones and siltstones, small lenses and few millimeter layers of siderite, chlorite-hydromica and chlorite-carbonate-hydromica with shells and fragments of microfauna of carbonate or cherty composition. Clastic sediments are of angular shape, unevenly distributed (10-40 %), represented by mica, silicites, feldspar and quartz. Authigenic minerals are represented by glauconite and pyrite, locally with siderite secretions. There are few fish scales and coal plant detritus.

Lithological structure of the Lower Maykop sequence of the Azov-Kerch region is characterized by a significant spatial-temporal variations

and the presence of reference horizons. In general, it is formed by 19 to 84 litmits of regional, zonal and local ranks with thickness of 20-80 m. They are represented by six lithmological complexes of four classification fields (fig. 3), the greatest spatial distribution of which (average 51 %) have the clay litmits (fields VI and VII).

Through development of psammite-aleuritic litmits is the almost traced in the south-western part of the region (the Northern-Vladyslavivska, Ust-Sal-

hirska and other areas). Here, the thickness of the Lower Maykop sequence is over 1000 meters and has a broad development of (more than 30 % of the section) mixed litmits (III) and aleurolitic (IV) classification fields.

An extensive development of clastogenic formations (55 %) is also found in central sections of the territories of the Kerch Peninsula (Krasnohirska, Kuibyshevskva, Fontanivska and other areas) and Near-Kerch shelf of the Black Sea (Subbotina area),

Table 1. Average calcium carbonate content in terrigenous rocks of Lower Maykop beds of Azov-Black Sea region

Area	Borehole №	Interval, m	Min. values	Max. values	Average, %	Species definitions	Rock
Alekseyevska	7	2875-3347	4,1	14,8	7,69	9	Arg., Sand.
Northern-Seleznivska	2	2794-3113	0,8	36,8	9,35	23	Sid., Arg., Sand.
Northern-Seleznivska	3	2803-3242	0	42,0	6,69	41	Sid., Arg., Sand.
Northern-Vladyslavivska	1	2529-3183	1,6	14,8	6,36	14	Sid., Arg., Sand.
Marfovskva	1	2993-3258	0,4	55,0	3,88	106	Arg., Sand.
Povorotna	1	2559-4012	0,8	26,0	4,01	27	Arg., Sand.
Bondarenkivskva	1	2225-3252	1,6	55,0	7,19	19	Sid., Arg., Sand.
Strilkova	8	970-1220	0	4,2	1,5	15	Arg., Sand.
Strilkova	7	1055-1656	0	2,05	0,95	3	Arg., Sand.
Peredova	1	1000-1008	0	0,4	0,15	8	Arg., Sand.
Ostrovskva	725	650-1230	0	9,6	2,2	17	Arg., Sand.
Ostrovskva	728	710-940	0	0,8	0	5	Arg., Sand.
Ostrovskva	726	350-1019	0	0	0	15	Arg., Sand.
Ostrovskva	727	872-1115	0	0	0	5	Arg., Sand.
Chongarska	705	336-1115	0	4,2	1,4	4	Arg., Sand.

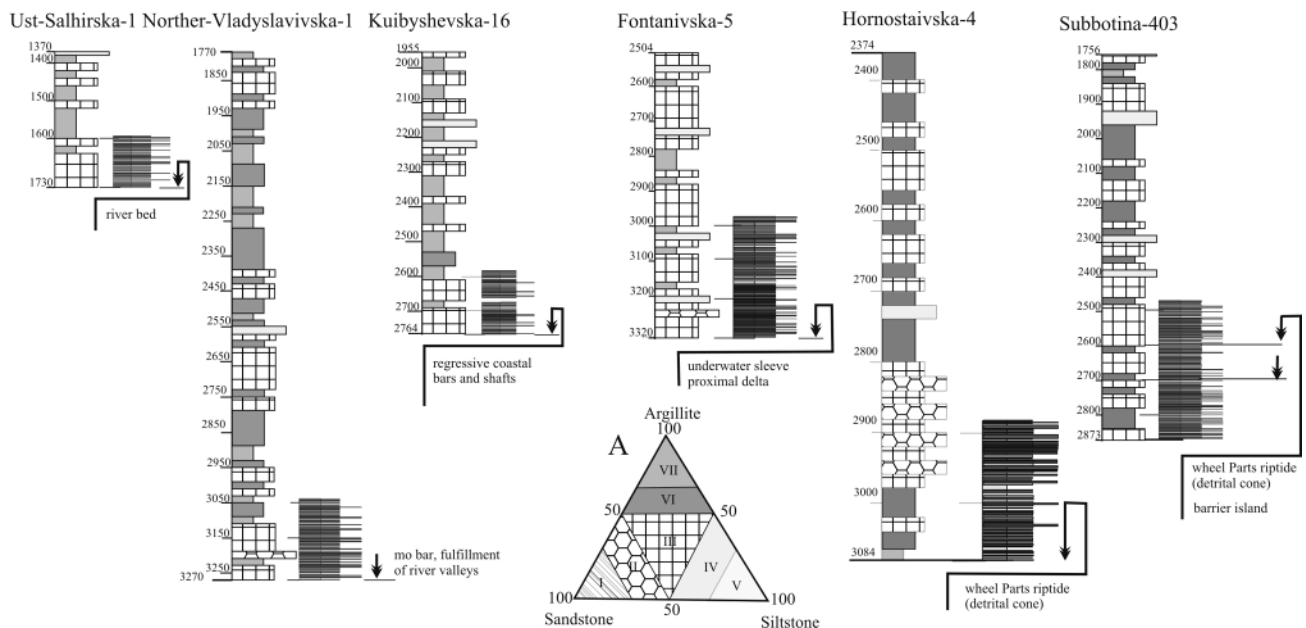


Fig. 3. Litmological sections of Lower Maykop and lithological structure of its basal layers with elements of sedimentological interpretation:

A – classification triangle. Legend see Fig. 2.

where they are represented by litmits of two classification fields: mixed (III, to 45 %) and aleurolitic (IV, 54 %). The lithmological structure of the latter one is formed by the siltstone layers (up to 8 of 0,1-1 m) with single sandstone (up to 0,2-1,5 m) and argillite (from 0,1-1,5 to 15 m) horizons.

In the rest area the content of psammite-aleurolitic litmits in section of the sequence does not exceed 25 %.

Quite clearly, that in the lithmological structure of the section of the Lower Maykop sequence the basal clastogenic pack of subregional plan takes the main part (fig. 3). It can be seen in the western and central parts of Indol-Kuban depression and the southern board of Sorokin depression. It is represented mainly by the formations of two fields: aleuritic (IV) and mixed (III), which can be replaced laterally. In the axial zones of consedimentary depressions these formations are split by the layers of clayey varieties (fields VI and VII). Sections of litmit is composed of interlaying of siltstones (thickness of individual horizons from 0,5-1 to 6 meters 25-50 % of section), sandstone (0,5-1,5 m 22-30 %) and argillite (from 0,5-1,5 to 7 m, 23-63 %) (fig. 3).

According to the result of the genetic interpretation of geophysical well logging of basal pack of the Lower Maykop sequence of the Azov-Kerch region, using the methodology (Muromtsev, 1983; Porebski, 1999), several major types of sedimentary environments can be distin-

guished (fig. 3). The following types of accumulation units have been diagnosed: river (mouth bars) or sea genesis (barrier islands, alongshore bars) (table 2).

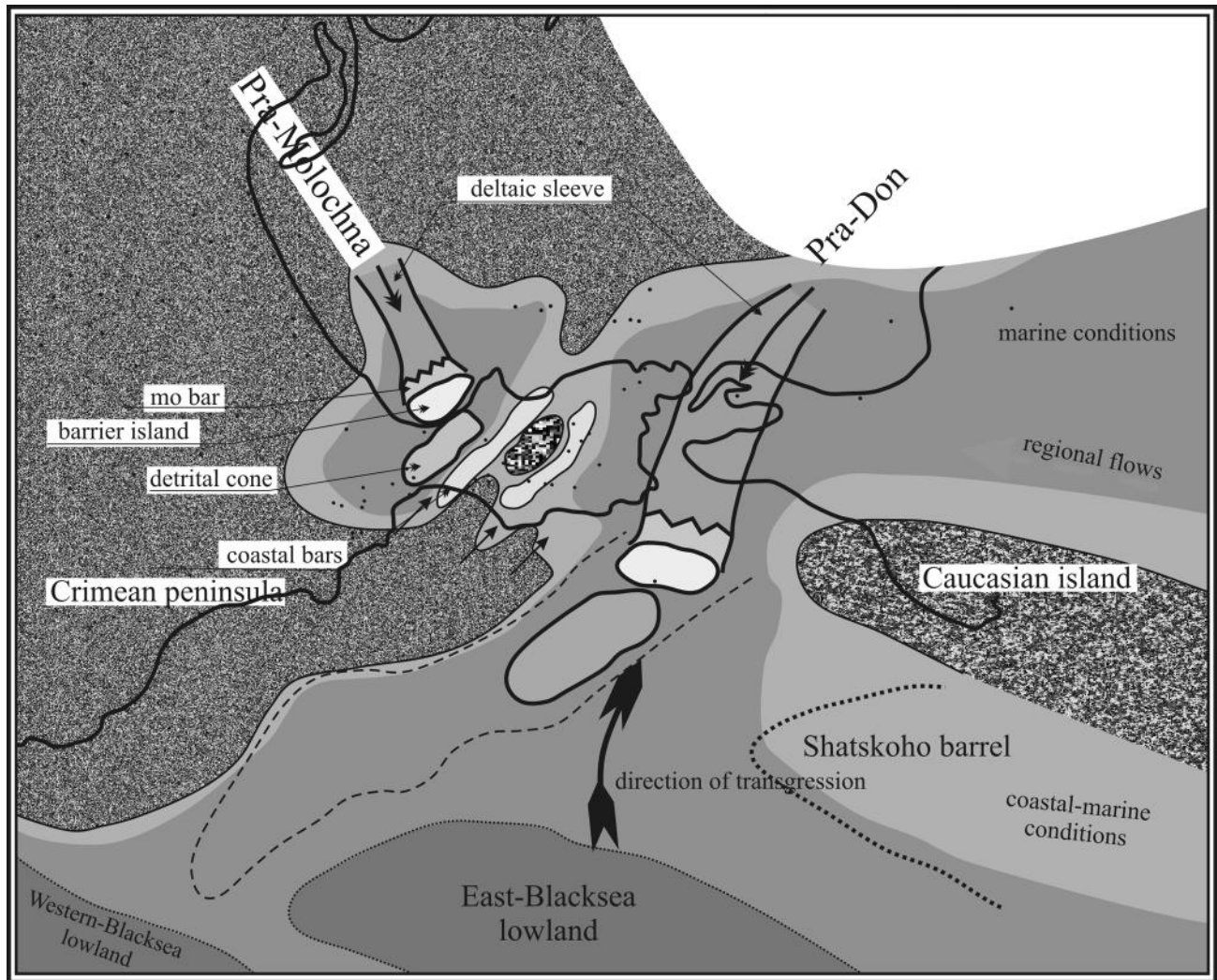
Based on the data above and taking into account the principle of facial balance (Selley, 1989) of dimensional stability of alluvial delta system structure: channel – bar of mouth – barrier island – detrital cone, a geological and paleoceanographic model of Azov-Kerch region for the beginning of early Maykop age was built (fig. 4).

**PALEOOCEANOGRAPHY OF EARLY MAYKOP**

The main factors that led to the nature of spacial-temporal occurrence of different by genesis Lower Maykop deposits of the region were the fore-Oligocene regression and the Early Oligocene transgression. First factor has led to the formation of complex erosial and tectonic topography with its branched system of river valleys, denudational highs and the local sedimentary basins, while the second one has caused the specific lateral diachronity of deposits. Taking into account the data on total paleoceanographic Oligocene situation in Paratethys basin (Kazmin, Tichonova, 2006) and the results obtained by us we can affirm that the depositional history that had begun in the range of modern Near-Kerch Black Sea shelf, was spread to the territories of Kerch peninsula, Plain Crimea and further to Near-Sivashsha.

Table 2. Diagnostics of near-shore marine sediments of basal horizons of Lower Maykop of Azov-Black Sea region

Area	Borehole №	Diagnostics of facies
Fontanivska	1	Underwater sleeve of proximal delta
Fontanivska	3	Shelf plain with passive hydrodynamics
Fontanivska	4	Underwater arm of proximal delta; estuarine bar
Fontanivska	5	Underwater arm of proximal delta
Fontanivska	6	Underwater arm of proximal delta
Fontanivska	7	River valley fill
Fontanivska	8	River valley fill
Ust-Salhirska	1	River valley fill
Dzhankoyska	18	Shelf
Korenkivska	1	Frontal parts of discontinuous currents
Hornostaiivska	4	Frontal parts of discontinuous currents
Feodosiyska	1	Regressive coastal bars and shafts
Kuibyshevsksa	16	Regressive coastal bars and shafts
Northern-Vladyslavivska	1	Mouth bar; River valley fill
Bulganaksksa	1	Underwater arm of proximal delta



**Fig. 4.** Paleoceanographic model of Azovo-Kerch region. Beginning of Early Maykop.

Analysing the nature of distribution of the sediments thickness, variations of aleuritic-psammitic components content and taking into consideration the data (Samarsky, 1983), we can affirm that local consedimentary paleohigh of the north-eastern stretch really existed on the territory of Kerch Peninsula. Moreover, the most pronounced was the Moshkarivsk-Selezniivska high zone, that significantly affected the dynamics of transport and distribution of terrigenous material.

Investigation of some terrigenous mineralogical indicators made it possible to allocate some source areas. Thus, the presence of Upper Eocene marl fragments, limestones, sandstones, quartziferous sandstones, pyroxene grains – indicators of a number of basic igneous rocks in the separate packs, indicates destruction of the southern (Crimean) paleoland. The latter is reflected in the established character of the spatial distribution of siltstones and sandstone deposits.

According to the mineralogical and petrographic studies (unsorted and unrounded clastic material, a significant number of heavy fraction minerals and their wide range, etc.), the pelitic and aleuritic-psammitic silts of the northern part of the basin were the products of hypergenic processing of Proterozoic highly metamorphosed rocks of the peneplained Feno-Sarmatian dryland. Moreover, in their spatial distribution, the main roles were played by alluvial systems of Pra-Don (East Region) and Pra-Molochna (west), that included the main by the volume of terrigenous material input into the sedimentary basin.

The distribution of terrigenous material in the range of the sedimentary basin has been determined by the influence of wave-cut, tidal and regional currents. As the sea level was growing the influence of tidal currents rised significantly, causing the destruction of accumulative alluvial deltaic bodies. In addition, according to (Zhyzhchenko, 1974), in the Early Maykop in the range of the

Western Caucasian basin the sublatitudinal flow existed that was directed from east to west, which caused the formation of alongshore bars and headlands. Obviously, that if the sea basin prograded, the current extended to the territories of the Azov-Kerch basin.

All this contributed to corresponding redistribution of clastic sediments along the Kerch paleoisland ridges with formation of sublatitudinal bands of accumulative bodies mainly the alongshore bars and headlands.

In the area of maximum subsidence (Vidnenska area) the accumulative delta complexes were formed (fig. 4).

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## CONCLUSIONS

At initial stages of the Early Oligocene evolution of Azov and Kerch segments of Paratethys the general nature of dynamics of transportation and distribution of terrigenous material within its ranges was defined by the spatial position, morphological and dynamic features of two powerful alluvial systems of Pra-Don and Pra-Molochna, existence of consedimentary highs (Moshkarivsko-Selezniivsk and other), which usually constituted the ridges of under- and above-water elevations and of wave-cut, tidal and regional currents, that led to redistribution of material with formation of accumulative bodies, mainly the alongshore bars and headlands.

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